

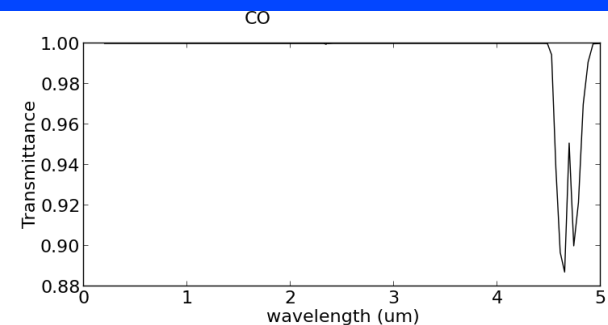
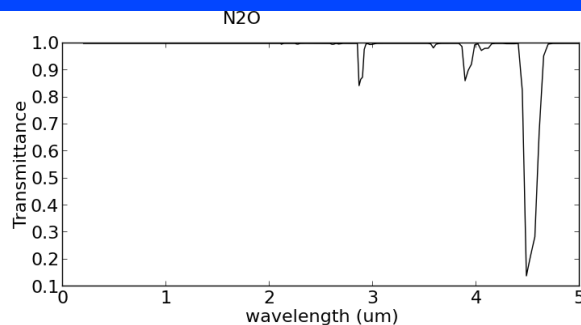
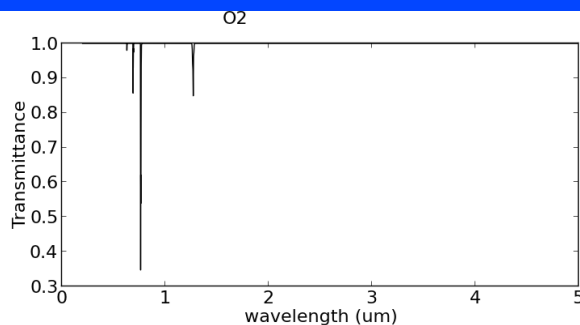
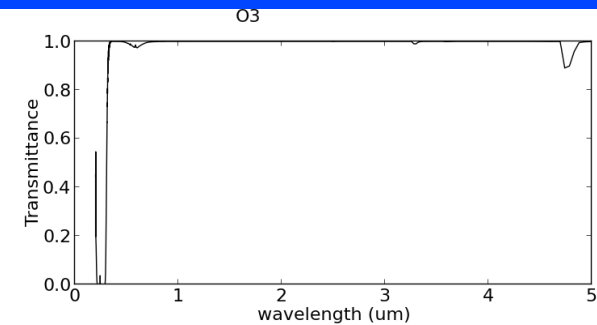
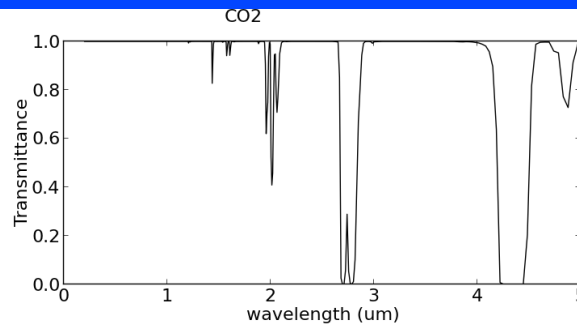
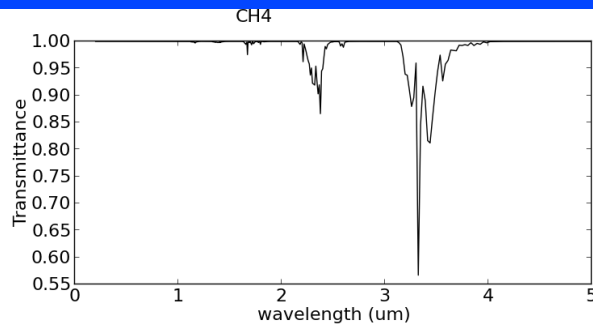
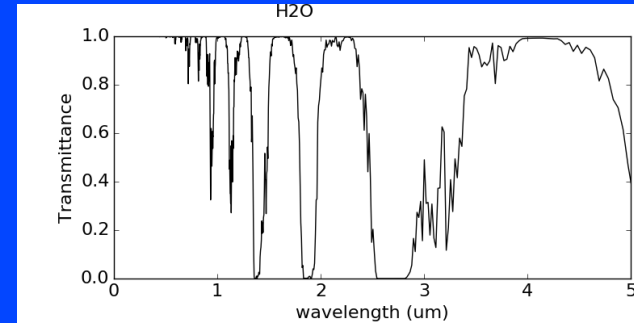
# Broadband radiative transfer simulations for un-filtering process of radiances from CERES

Lusheng Liang<sup>1</sup> and Wenying Su<sup>2</sup>

1-SSAI; 2-NASA LaRc

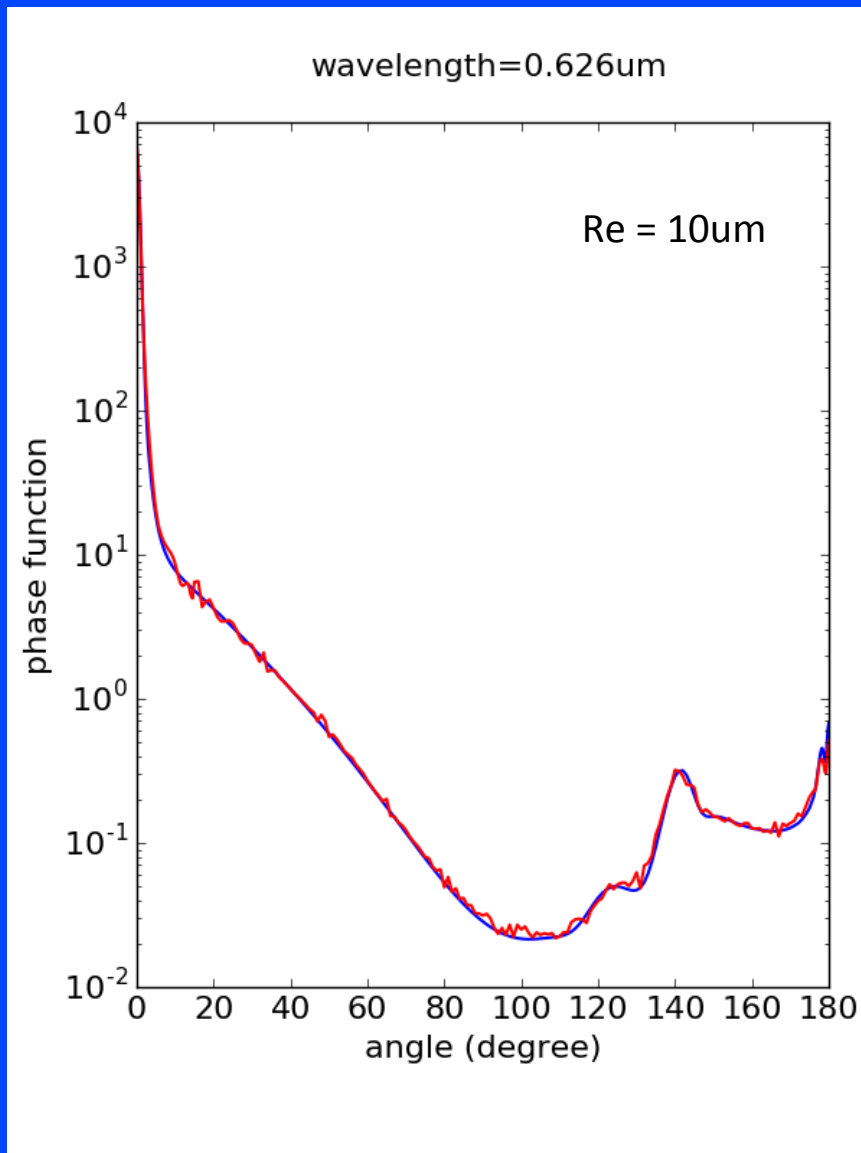
# 1. Absorption of atmospheric gases in shortwave

- Absorption properties of 7 major gases ( $\text{H}_2\text{O}$ ,  $\text{O}_3$ ,  $\text{O}_2$ ,  $\text{CH}_4$ ,  $\text{CO}_2$ ,  $\text{CO}$  and  $\text{NO}$ ) in atmosphere are all derived from line parameter database (aer\_v\_3.2 based on HITRAN 2008) of LBLRTM code;
- The previous version of broadband radiative transfer code considered absorption properties of  $\text{H}_2\text{O}$ , and  $\text{O}_2$  derived from LBLRTM line parameter database and the that of  $\text{O}_3$  was derived from WMO's Atmospheric Ozone 1985.



for the U.S. 1962 standard atmospheric profile

## 2.1 Cloud optical properties--water clouds



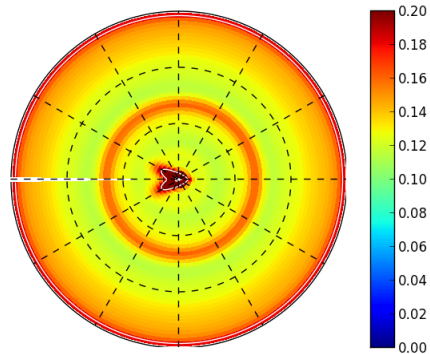
(1) Ping Yang's water cloud optical properties are used in the radiative transfer simulations

(2) Testing phase function (blue curve):

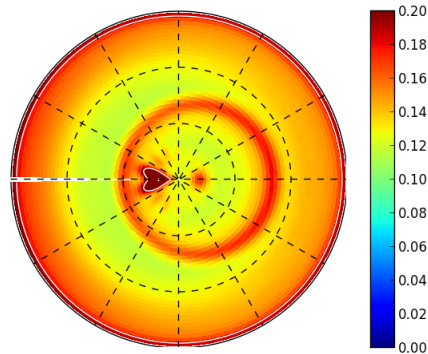
- log-normal distribution with the standard deviation of 0.35
- the range of the particle size distribution is from 0.001 um to 100 um with an interval of 0.001 um

# BRF for water clouds based on the cloud optical properties

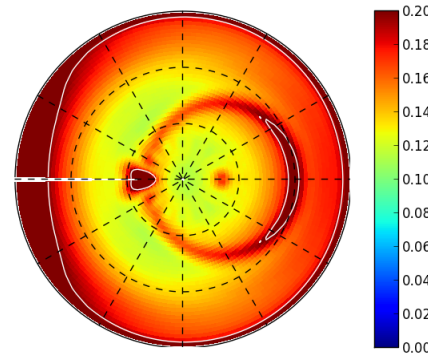
SZA=1



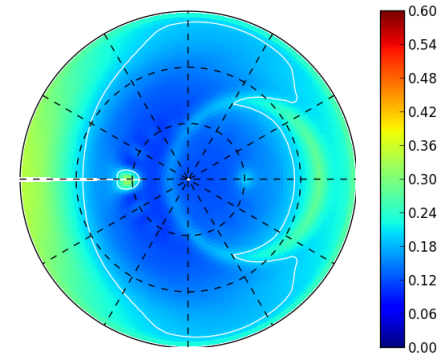
SZA=10



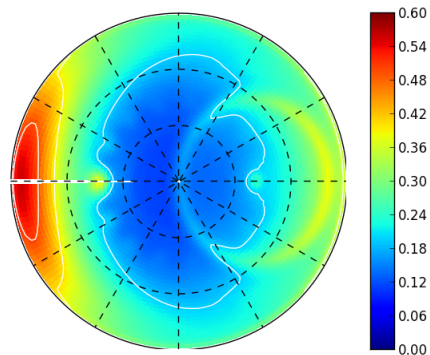
SZA=20



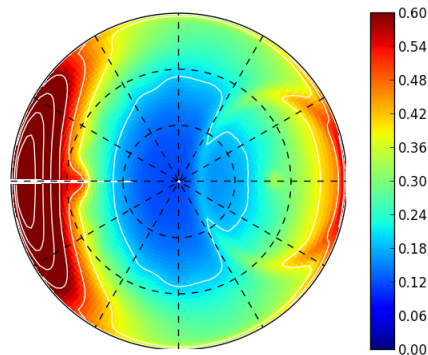
SZA=30



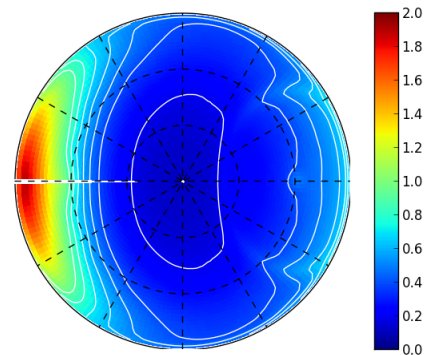
SZA=40



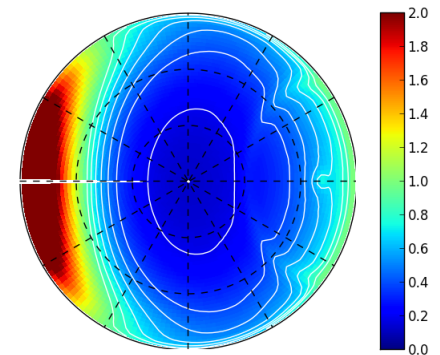
SZA=50



SZA=60



SZA=70



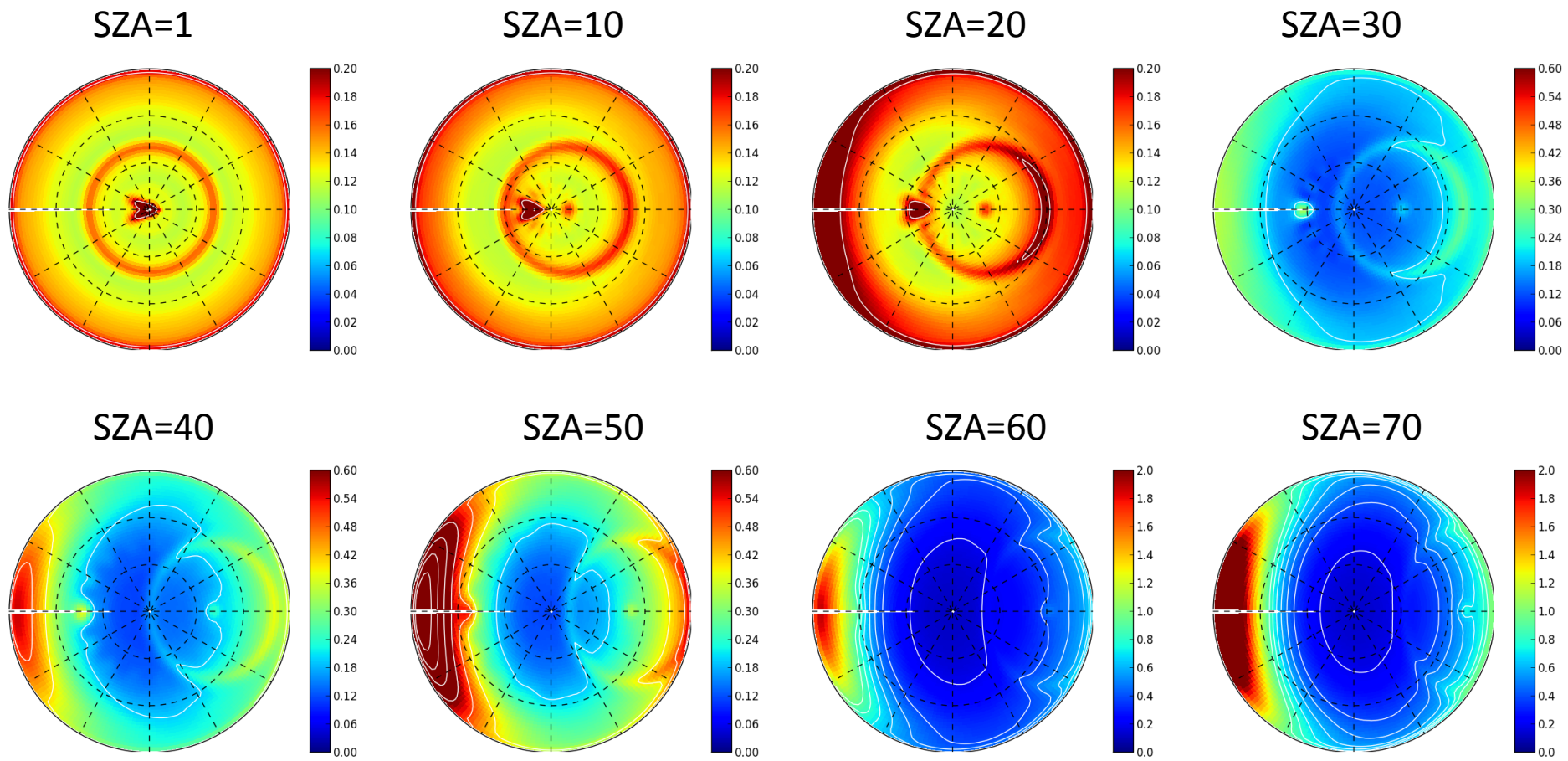
wavelength ranging from 0.626-0.714 $\mu$ m  
cloud optical depth=2  
ocean surface wind speed=0 m/s  
effective radius of clouds = 10 $\mu$ m

## 2.2 Cloud optical properties—ice clouds

- (1) The particle effective diameter is large relative to the wavelength in the shortwave;
- (2) Very large peaks appear in the forward scattering directions close to 0 degree in the phase functions, which requires a lot of legendre polynomials to represent phase function in radiative transfer simulation and computations are too expensive;
- (3) Phase functions need to be modified;
- (4) Radiative transfer codes like DISORT are unable to produce the correct radiances with limited terms of legendre polynomials to represent phase function even with delta-M truncation is turned on for ice clouds.

# Handling high peaks of ice cloud phase functions

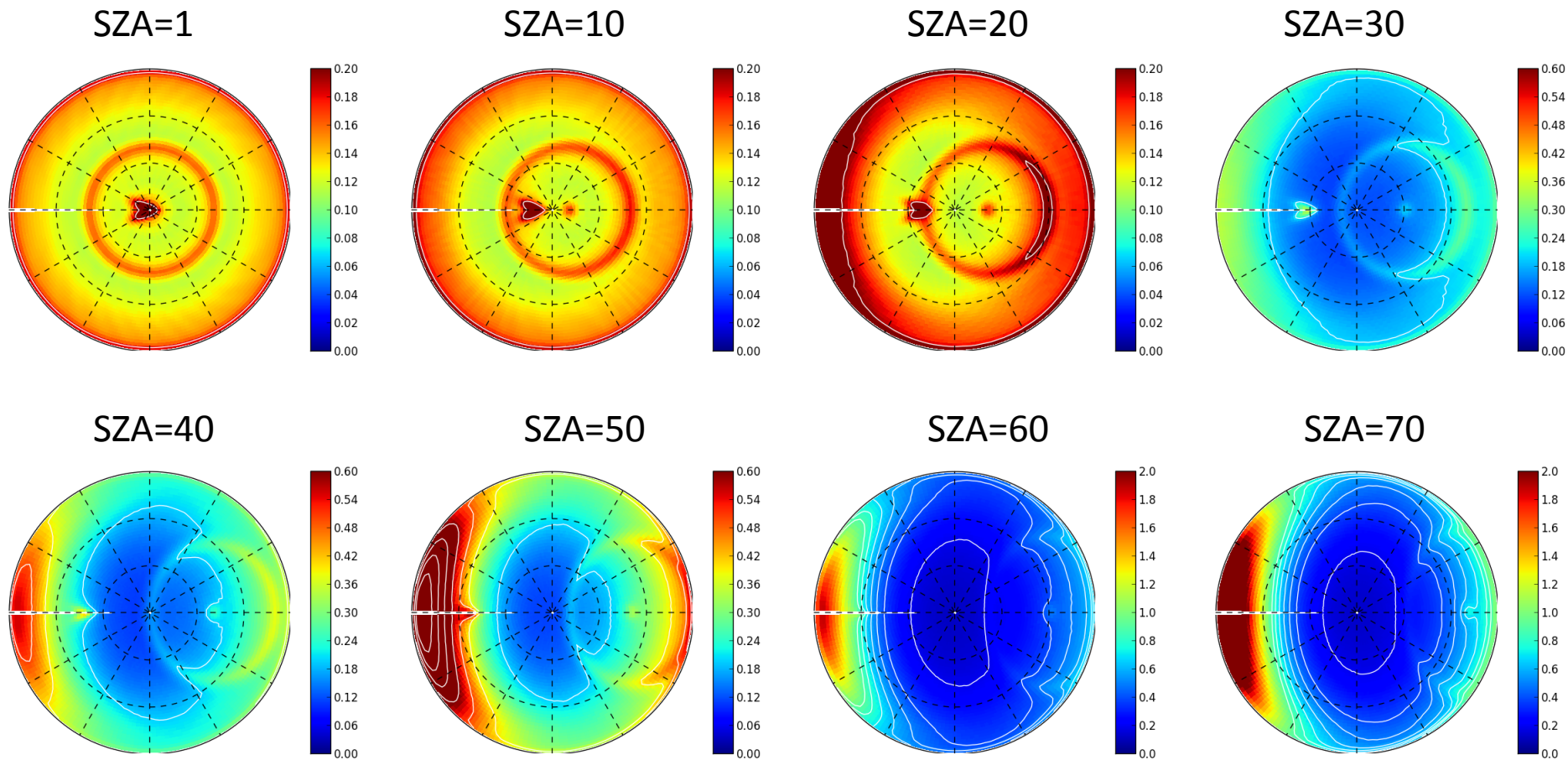
Reference BRF based on water cloud with  $Re=10\mu m$



wavelength ranging from 0.626-0.714um  
water cloud optical depth=2  
ocean surface wind speed=0 m/s  
effective radius of clouds = 10um

# Handling high peaks of ice cloud phase functions

BRF calculated based on a geometry-cut algorithm (Iwabuchi and Suzuki, 2009)

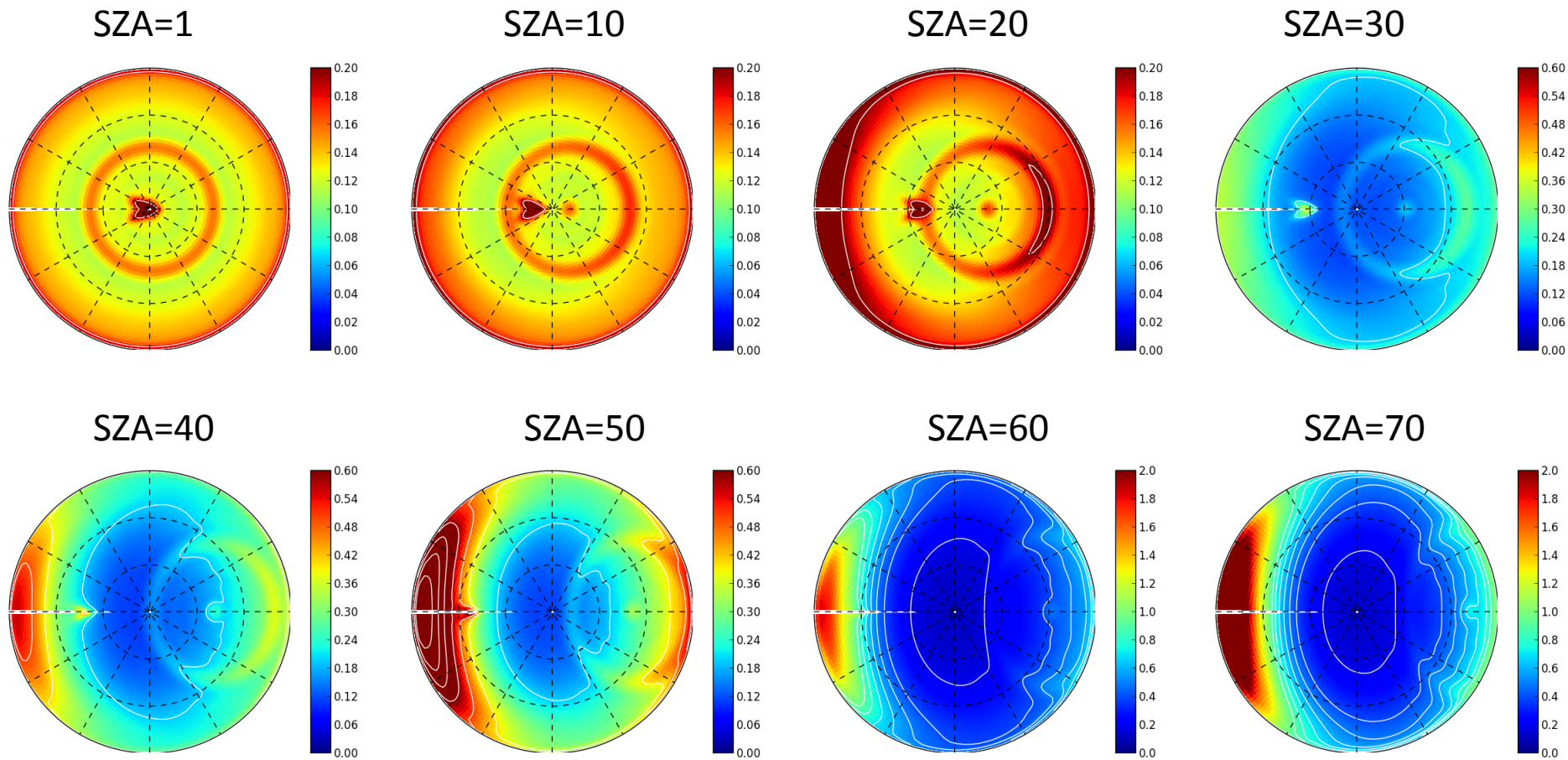


wavelength ranging from 0.626-0.714 $\mu$ m  
water cloud optical depth=2  
ocean surface wind speed=0 m/s  
effective radius of clouds = 10 $\mu$ m



# Handling high peaks of ice cloud phase functions

BRF calculated based on the delta-fit algorithm (Hu et al., 2000)



wavelength ranging from 0.626-0.714 $\mu$ m  
water cloud optical depth=2  
ocean surface wind speed=0 m/s  
effective radius of clouds = 10 $\mu$ m



### 3. Land surface Bidirectional Reflectance Distribution Functions

At a given wavelength, the BRDF is calculated based on Ross-Li BRDF model

$$R = P_0 + P_1 K_1(\theta_0, \theta, \phi) + P_2 K_2(\theta_0, \theta, \phi)$$

where  $K_1(\theta_0, \theta, \phi)$  and  $K_2(\theta_0, \theta, \phi)$  are two kernels, and  $P_0$ ,  $P_1$  and  $P_2$  are the BRDF parameters.

For every IGBP surface type:

- At wavelengths of MODIS 7 land bands (bands 1-7), BRDF parameters were averaged from MODIS 10-year observations (MCD43C1)
- The parameters at other wavelengths between 0.47 and 2.1um are calculated with spline interpolations
- The parameters at the wavelengths  $< 0.47$  um or  $> 2.1$  um are scaled along with the spectral reflectance from JPL spectral library 2.0.

## 4. Broad-band radiative transfer simulations over land

- Shortwave radiations are computed at a spectral resolution of  $20\text{cm}^{-1}$  from  $0.2\text{ }\mu\text{m}$  to  $5.0\text{ }\mu\text{m}$  with a radiative transfer code based on DISORT.
- 16 streams for calculations containing water clouds or aerosols and 32 streams for calculations containing ice clouds.
- Broad-band radiative transfer simulations are performed for 7 IGBP surface types
  - 1: Evergreen Needleleaf Forest
  - 3: Deciduous Needleleaf Forest
  - 4: Deciduous Broadleaf Forest
  - 5: Mixed Forest
  - 10: Grasslands
  - 11: Permanent Wetlands
  - 16: Bare Soil and Rocks
- 6 bins of cosine of SZA (1.0, 0.85, 0.7, 0.55, 0.25, and 0.1)
- 9 bins of VZA (5, 15, 25, 35, 45, 55, 65, 75, and 85)
- 9 bins of RZA (5, 25, 45, 65, 85, 105, 125, 145, 165)
- 4 bins of aerosol optical depths (0, 0.05, 0.14 and 0.30, OPAC dust model for IGBP=16 and continental average aerosol model for other IGBP surface types)
- Ping Yang's water and ice cloud optical properties
- US 1976 standard atmosphere model

## 5. SW Radiance Unfiltering processes

1. Unfiltered, reflected radiance is calculated from filtered, reflected radiances as:

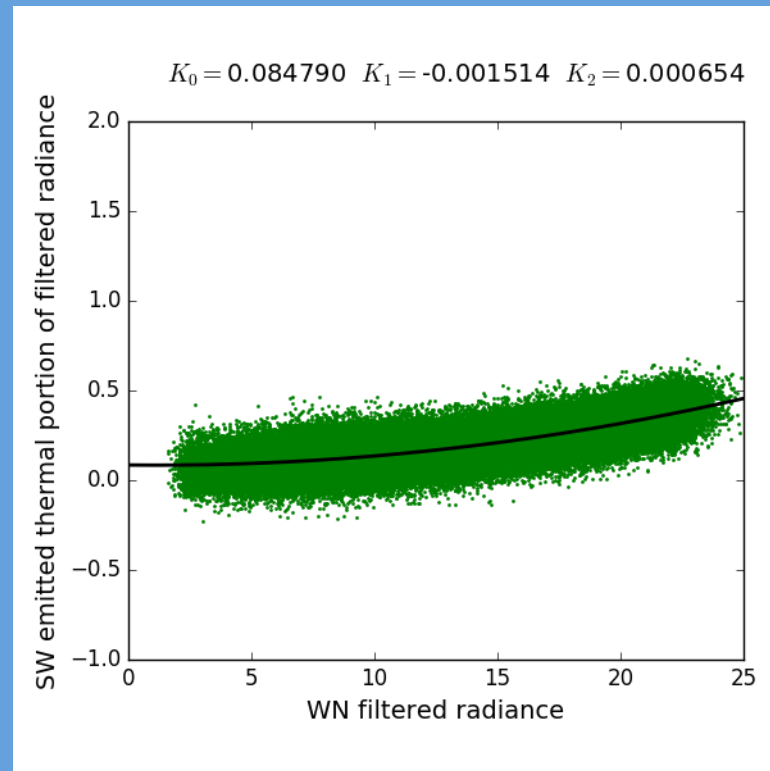
$$m_u^{sw} = a_0 + a_1(m_f^{sw_r}) + a_2(m_r^{sw_r})^2$$

2. The filtered, reflected radiance is calculated as:

$$m_f^{sw_r} = m_f^{sw} - m_f^{sw_e}$$

3. The emitted thermal portion of the filtered radiance is calculated as:

$$m_f^{sw_e} = K_0 + K_1(m_f^{wn}) + K_2(m_f^{wn})^2$$



The data are extracted from the first day in each month in 2001 of Ed4 night time Terra FM1. Required SZA > 95 degrees to exclude the impact of twilight.

## 6. CERES SW radiance unfiltering coefficients as constructed from simulated reflected radiances

(1). Calculate unfiltered broadband radiances (no emitted)

$$m_u^{sw} = \int_0^{\infty} I_{\lambda}^r d\lambda$$

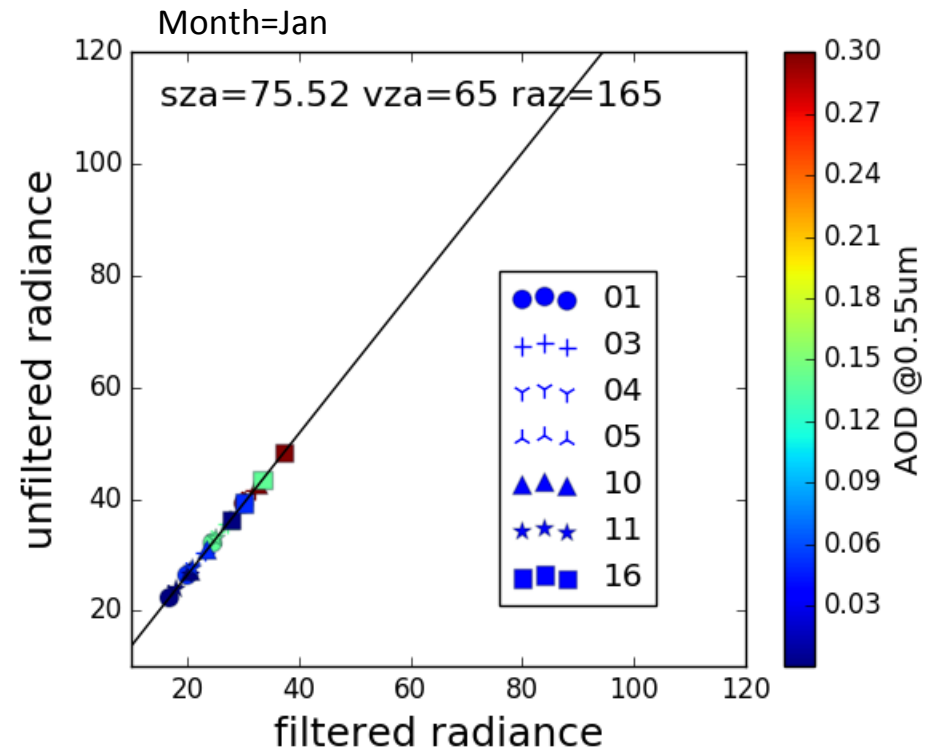
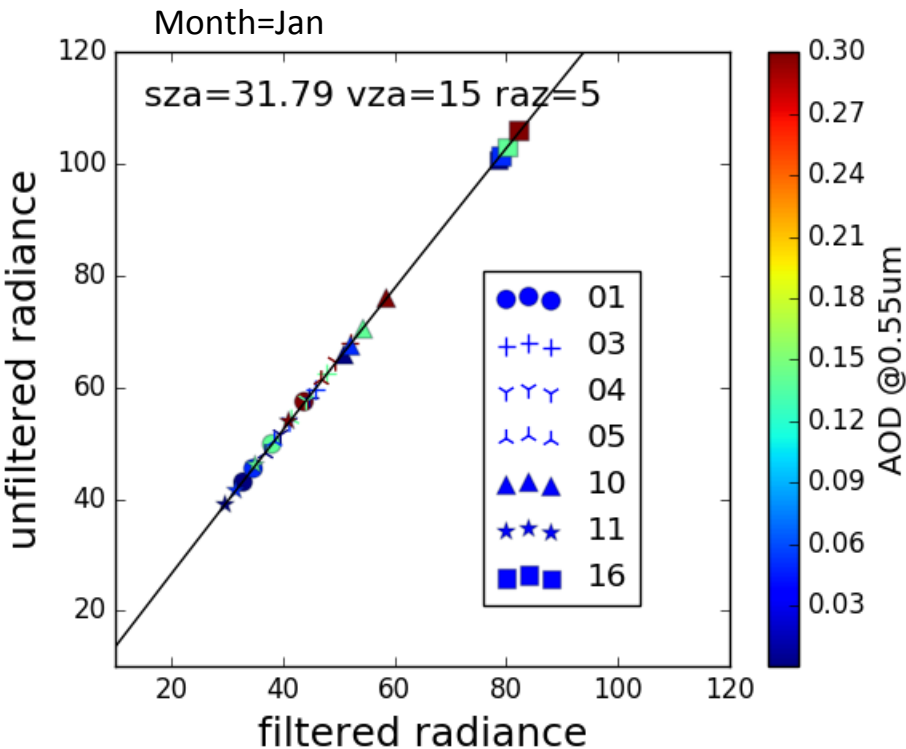
(2). Apply CERES spectral response functions to calculate filtered broadband radiances

$$m_f^{sw} = \int_0^{\infty} S_{\lambda}^j I_{\lambda}^r d\lambda$$

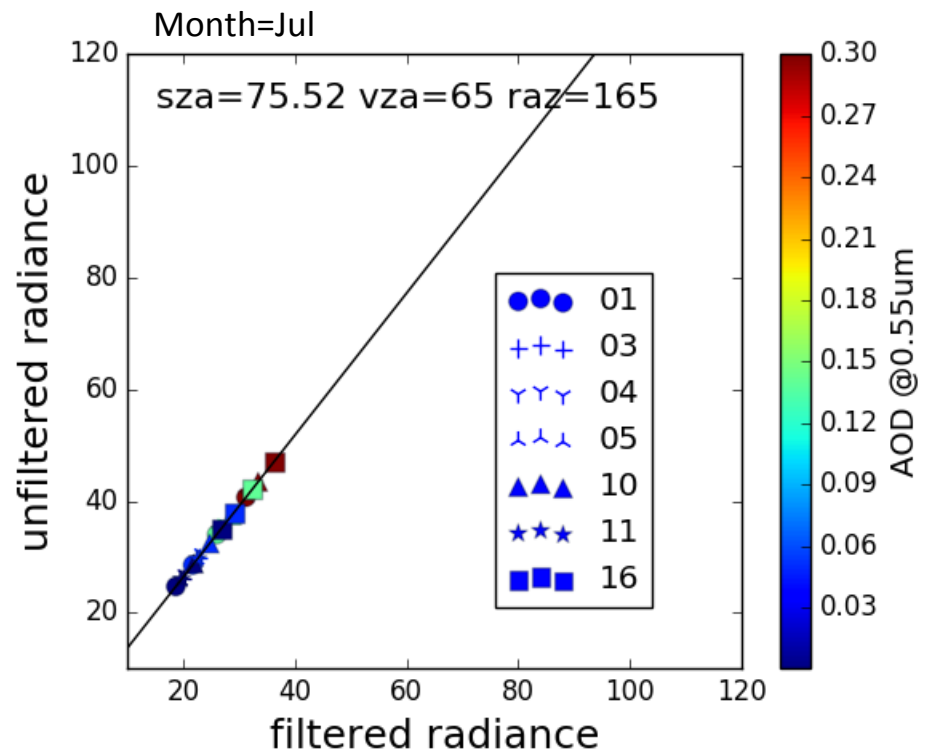
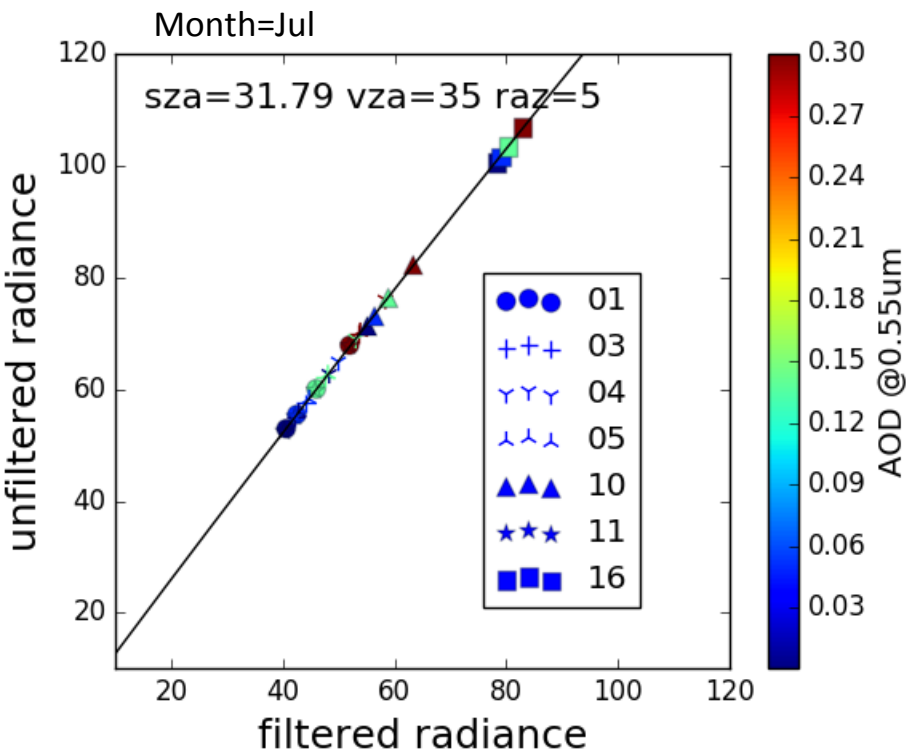
(3). Convert filtered radiances to unfiltered radiances by

$$m_u^{sw} = a_0 + a_1(m_f^{sw_r}) + a_2(m_f^{sw_r})^2$$

(4). Calculate unfiltering coefficients ( $a_0$ ,  $a_1$  and  $a_2$ ) in every SZA-VZA-RAZ bin. Within each bin, as for the currently available simulations, 28 pairs of data are used (7 IGBP types X 4 AOD).

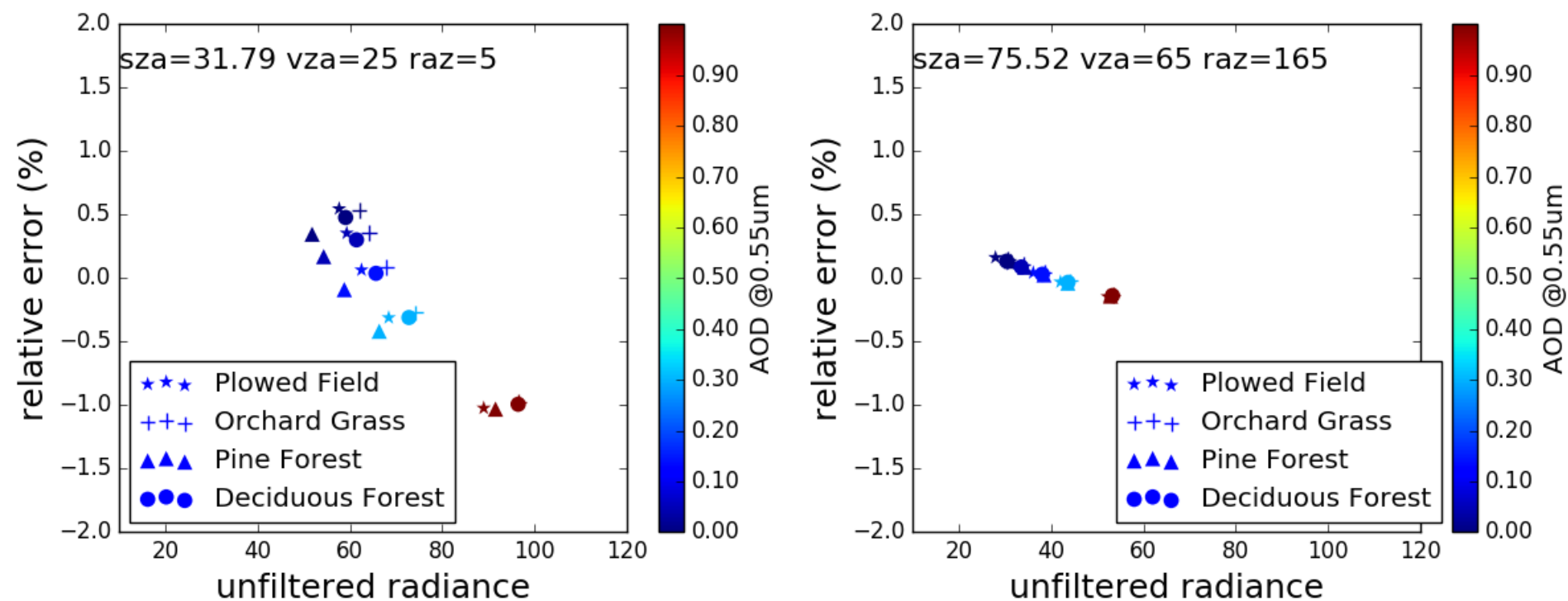


The differences in coefficients (as shown in the fitting curve) between January and July are small.



## 7. Errors in unfiltering processes over land

- (1) Calculate broadband radiances (both filtered and unfiltered) with alternative BRDF models (Roujean et al, 1992). The surface models of plowed field, orchard grass, pine forest and deciduous forest are simulated; aerosol optical depths are 0, 0.05, 0.14, 0.30 and 1.0 of OPAC continental average model.
- (2) Take the filtered radiances in (1) and apply unfiltering coefficients to get the unfiltered radiances.
- (3) compare the unfiltered radiances in (2) to the unfiltered radiances in (1).





# Future work

- (1) Complete clear-sky unfiltering processes over land based on the simulations for all IGBP surface types
- (2) Complete unfiltering processes for snow and sea ice
- (3) Investigate the unfiltering processes for footprints partially covered by fresh snow or sea ice
- (4) complete unfiltering processes over ocean
- (5) Repeat (1) to (4) for cloudy conditions
- (6) Repeat (1) to (5) in the longwave range